

Vegetation history around Yongneup moor at Mt. Daeamsan, Korea

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Abstract

A pollen analytical study of sediment sequences collected from Yongneup moor (sampling point: 38°12'57.4" N, 120°7'30.2" E) was conducted to understand the vegetation history in the mountainous region of the central Korean peninsula. Carbon dating was carried out to measure five successive samples obtained from the bottom at a depth of 180 cm to the surface. The Yongneup moor sediment revealed four main local pollen zones; that is, four past vegetation phases as follows: Local pollen zone I: *Quercus-Pinus* zone; estimated age, 5,900-4,800 calibrated years (cal) before present (BP); vegetation type, cool-temperate central/montane deciduous broad-leaved forest. Local pollen zone II: *Pinus-Abies-Quercus* zone; estimated age, 4,800-3,400 cal BP; vegetation type, cool-temperate northern/alti-montane mixed coniferous and deciduous broad-leaved forest. Local pollen zone III: *Quercus-Pinus-Abies* zone; estimated age, 3,400-400 cal BP; vegetation type: cool-temperate central/montane deciduous broad-leaved forest. Local pollen zone IV: *Pinus-Quercus* zone; estimated age, 400-present cal BP; vegetation type, cool-temperate central/montane mixed deciduous broad-leaved and coniferous forest. It was confirmed that subalpine coniferous forests had expanded to the mountainous region of the central Korean peninsula during the period from 4,800-3,400 cal BP and thereafter deciduous forests dominated by *Q. mongolica* were established. Notably, secondary forests dominated by *P. densiflora* developed in the lower part of the mountainous region of the central Korean peninsula about 400 cal BP due to human interference.

Key words: cool-temperate forest, Holocene, pollen analysis, vegetation history, Yongneup moor

INTRODUCTION

Environmental change is a continuous process that is well recorded in only a few ecosystems. Among these are high moors (bogs), which are remarkable archives of long-term natural changes, because they document vegetation history within peat that accumulates almost continuously over several millennia (Charman 2002, Robichaud and Bégin 2009). The development of bogs is influenced in a complex combination by autogenic processes (peat build-up) and allogenic forces such as climate, disturbances, and topography (Tolonen 1987, Korhola 1992, Kuhry et al. 1993, Charman 2002).

Yongneup moor is developing at the northwestern

slope of Mt. Daeamsan (1,304 m) at the boundary between Inje-gun county and Yanggu-gun county, Gangwon-do province, Korea. Yongneup was first classified as a high moor (German classification) or a raised bog (England classification) by the demilitarized zone (DMZ) survey team in 1967, and, thereafter, much attention has been paid to the Yongneup moor because it is a unique raised bog in Korea (Kang et al. 2010).

Pollen analysis is the principal technique used to document and reconstruct past vegetation and environments of the local region around raised bogs (Birks and Birks 1980). Since Yamazaki (1940) reported the first pol-

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len analysis of Saeseok Pyeongjeon (a wet grassland) at Mt. Jirisan located in the southern part of Korea, a period followed of over 30 years in which pollen analysis studies and reconstructions of vegetation history were not conducted in Korea. By the end of the 1970s, studies using accurate ^{14}C -dating were conducted on sediments in the east-central lowlands and the sub-mountain areas including Lake Youngrangho on the Korean peninsula (Jo 1979, Yasuda et al. 1978, Yoon 1998, Fujiki and Yasuda 2004, Park 2005).

However, studies on changes in past vegetation obtained from mountainous or subalpine wetland sediments were not performed except for those at Yongneup moor, which developed in the mountainous area of the east-central inland of the Korean peninsula. Therefore, it was assumed that Yongneup bog would be a suitable site to reconstruct the development and changes in past local vegetation. Previous pollen analytical studies on this moor (Kang 1980, Chang et al. 1987) were limited in reconstructing past vegetation changes, because their radiocarbon dating results were insufficient and they collected a small number of peat samples.

In the present study, we attempted to reconstruct changes in the local vegetation of the mountainous area of the central Korean peninsula using pollen analysis in parallel with radiocarbon dating of a number of samples collected from the peat sediments of Yongneup moor known as a unique high moor (raised bog) in Korea. The results obtained from this study will provide information for understanding the past local vegetation around Mt. Daeamsan.

MATERIALS AND METHODS

Study site

Yongneup moor is a depressive wetland with an area of 7,490 m² on the northwestern slope of Mt. Daeamsan (Fig. 1). The surrounding area (1.06 km²), including Yongneup, was first designated as an Ecosystem Conservation Area in 1989 but an expanded area of 1.36 km², including two other nearby wetlands, were considered an Ecosystem Conservation Area in 2010.

Furthermore, the Yongneup moor itself was registered as a Ramsar Convention Wetland on March 28, 1997 for the first time in Korea and Yongneup was designated as a Wetland Conservation Area in 1999. Additionally, nearby mountains such as Mt. Dosolsan (1,478 m), Daewoosan (1,179 m), Baekseoksan (1,142 m), and Samyeongsan

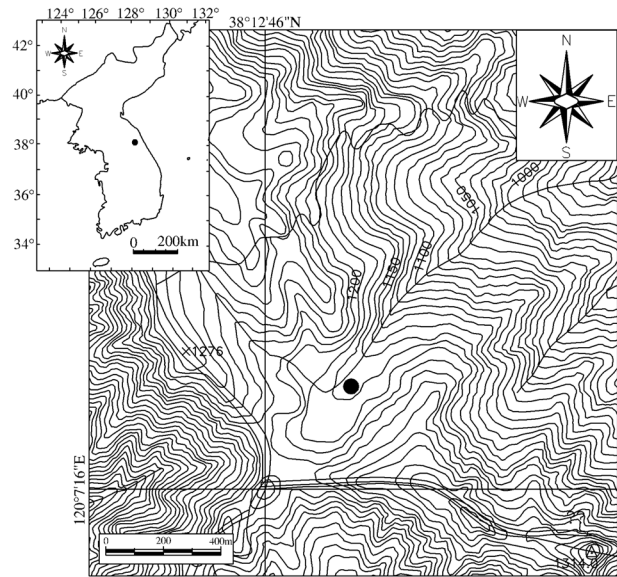


Fig. 1. Map showing the sampling site (●) of the Yongneup moor at Mt. Daeamsan, Gangwon-do province, Korea.

(1,198 m) including Mt. Daeamsan were designated as Nature Reserves in 1973.

Vegetation around Yongneup moor

Actual vegetation on the mountains around Yongneup moor is composed of deciduous broad-leaved forest dominated by *Quercus mongolica*. Besides *Q. mongolica*, *Abies nephrolepis* and *Pinus koraiensis* are mainly scattered on mountains over 1,000 m and on rocky ridges near the summit of Mt. Daeamsan, respectively.

In contrast, *Pinus densiflora* is rarely observed on mountains higher than 1,200 m, but appear as mixed forests with *Q. mongolica* with decreases in elevation. At the foot of the mountain, mixed forests of *P. densiflora* and *Quercus variabilis* occur. *P. densiflora* is the dominant tree species in residence areas (Min et al. 2000, Ministry of Environment of Korea 2007).

Climate

To understand the climatic characteristics around Yongneup moor, we constructed a climate-diagram (Walter et al. 1975) using meteorological information measured at Inje Meteorological Station during the 30 years from 1979-2008 (Fig. 2).

The average temperature was 10.0°C and annual precipitation was 1,187.5 mm. The mean temperature in August (warmest month) is 23.2°C and that in Janu-

ary (coldest month) is -5.1°C . Approximately 60% of all precipitation occurs from June to August. A period of greater than 100 mm of monthly precipitation occurs for 4 months from June to September. Such a precipitation pattern is similar to other regions on the Korean peninsula. In particular, a severe dry month, limiting plant growth, does not occur at Yongneup moor.

Pollen analysis and radiocarbon dating

A pollen analysis is a tool to reconstruct vegetation history from a particular location over time. Sediment sample cores for the present study were collected using a Hiller type peat-borer at $38^{\circ}12'57.4''\text{ N}$, $120^{\circ}7'30.2''\text{ E}$ (Fig. 1). As shown in Fig. 3 and Table 1, five successive samples for ^{14}C -dating were obtained from the bottom, at a depth of 180 cm, to the surface and were transported to the AMS facility at Seoul National University. Radiocarbon years were converted to calendar years using the CalPal program available online (Danzeglocke et al. 2008). The 61 sediment samples for pollen analysis were also sub-sampled in parallel with a description of sediment properties using revised soil color charts (Oyama and Takehara 1997).

Standard procedures, including 10% KOH treatment, heavy-liquid separation with ZnCl_2 (specific gravity 1.68-1.70), and subsequent acetolysis (Erdtman 1943, 1960, Fægri and Iversen, 1989) were used to extract pollen from the sediment samples. Pollen and spores mounted in glycerin jelly were counted under a light microscope at magnifications of $\times 400$ -1,000.

Identification of pollen and spores was aided using previously published studies (Wodehouse 1935, Erdtman 1986, Fægri and Iversen 1989) and by comparing the pollen and spores to reference pollen samples collected and treated at the University of Ulsan. More than 300 pollen grains and spores were counted based on the amount of arboreal pollen (AP) per slide. A pollen diagram was created using relative pollen frequency (RPF) based on the AP.

RESULTS

Identified pollen and spores were classified into AP (trees) and non-arboreal pollen (NAP, shrubs, herbs, and spores). The scientific names used are the families or genera of the pollen corresponding to the plant.

AP was as follows: *Abies*, *Acer*, *Betula*, *Carpinus*, *Castanea*, *Celtis*, *Fraxinus*, *Juglans*, *Picea*, *Pinus*, *Platycarya*,

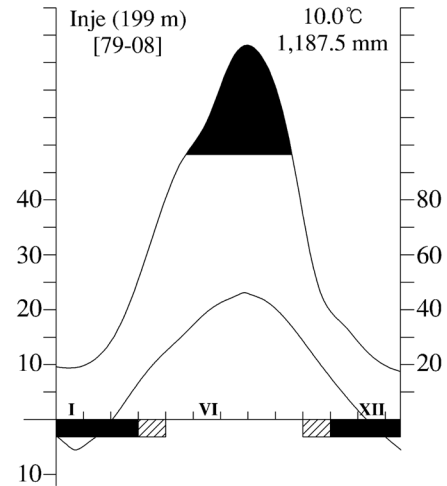


Fig. 2. Climate diagram for Inje-gun county, Gangwon-do province, Korea.

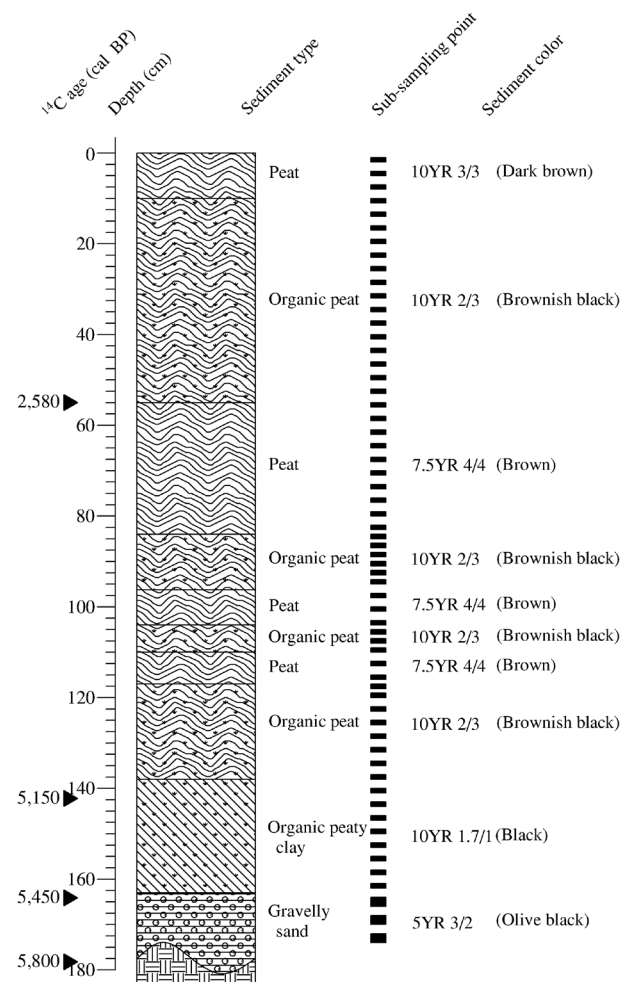


Fig. 3. Vertical facies succession of the sediment core collected from the Yongneup moor at Mt. Daeamsan, Gangwon-do province, Korea. Cal BP, calibrated years before present.

Pterocarya, Quercus, Tilia, Tsuga, and Ulmus + Zelkova.

NAP was as follows: *Alnus*, Araliaceae, Celastraceae, *Corylus*, *Lespedeza*, *Ligustrum*, *Lonicera*, *Meliosma*, *Rhododendron*, *Rhus*, *Salix*, *Staphylea*, *Symplocos*, *Weigela*, *Aconitum*, *Artemisia*, *Astilbe*, *Bistorta*, Caryophyllaceae, Chenopodiaceae, *Clematis* + *Caltha*, Cruciferae, Cucurbitaceae, Cyperaceae, *Drosera*, *Eriocaulon*, *Geranium*, Gramineae, *Haloragis*, *Impatiens*, *Iris*, Juncaceae, *Leonurus* type, Liliaceae, *Lysimachia*, *Lythrum*, *Menyanthes*, Onagraceae, other Compositae, other Labiatae, other Rosaceae, *Parnassia*, *Patrinia*, *Persicaria*, *Plantago*, *Polygonum*, *Ranunculus*, *Sagittaria*, *Sanguisorba*, *Scabiosa*, *Sedum*, *Thalictrum*, *Typha*, Umbelliferae, *Vicia*, monolete spores, and trilete spore.

Physical characteristics of sediment

The vertical facies of the sediment core collected from the Yongneup moor are presented in Fig. 3. Of the entire 180 cm, the upper layer from a depth of 138 cm to the surface was mainly composed of peat or organic peat.

The layer between 180-163 cm was gravelly sand containing some clay, and the color of these layers was olive black due to the small quantity of organic matter. The layer between 163-138 cm was organic peaty clay containing indistinguishable organic remains and distinguishable plant debris with the naked eye. The color of these layers was black. The layer between 138-10 cm alternated repeatedly with brownish and humified peat, indicating the relative difference in the humification of peat. The layer from 10 cm to the surface was composed of dark brown peat, suggesting that this upper layer was affected and disturbed by the peat sampling procedure because of its high water content.

From the physical features described above, soil components from the vicinity of the slope at Yongneup moor did not flow from the 138 cm depth layer to the surface. Accordingly, it was inferred that the peat sediment of

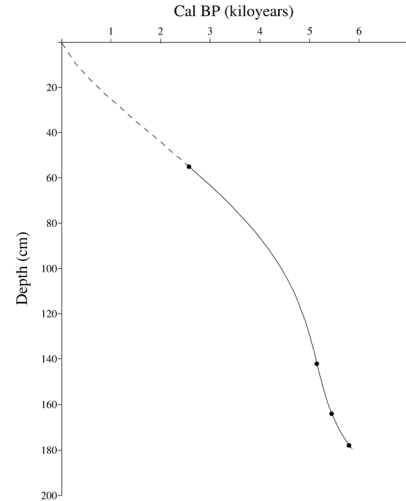


Fig. 4. Age-depth curve based on the accelerator mass spectrometry radiocarbon dating results obtained from sediment samples of Yongneup moor at Mt. Daeamsan, Gangwon-do province, Korea. Cal BP, calibrated years before present.

these upper layers was only composed of autochthonous plant remains, including air-born pollen and spores, as well as allochthonous dust and aerial deposition.

¹⁴C-dating

It was confirmed by ¹⁴C-dating that the Yongneup moor was initiated by the formation of a concave wetland approximately 5,900 calibrated years (cal) before present (BP) (Table 1). An age-depth curve was plotted in Fig. 4 to estimate the sediment ages using the ¹⁴C-dating results.

Pollen analysis

To facilitate discussion of the pollen record, the pollen diagram was subdivided into four local pollen zones (LPZ) based on the percentage of the principal pollen and spores (Fig. 5)

Table 1. AMS radiocarbon ages of sediment samples obtained from the Yongneup moor on Mt. Daeamsan, Gangwon-do province, Korea

Sample No.	Depth level (cm)	Dated material	Lab. No.	Age in ¹⁴ C BP	Age in cal BC/AD
Yongneup-site 6-1	54-56	Sediment	SNU08-722 [†]	2,510 ± 50	630 BC
Yongneup-site 6-2 [*]	88-90	Sediment	SNU08-723	1,800 ± 50	220 AD
Yongneup-site 6-3	141-143	Sediment	SNU08-724	4,480 ± 50	3,200 BC
Yongneup-site 6-4	163-165	Sediment	SNU08-725	4,740 ± 50	3,500 BC
Yongneup-site 6-5	177-179	Sediment	SNU08-726	5,060 ± 50	3,850 BC

AMS, accelerator mass spectrometry; BP before present; Cal, calibrated years; BC, before Christ; AD, Anno Domini.

^{*}Eliminated to construct the age-depth curve.

[†]SNU 08: code number for the AMS Center of Seoul National University.

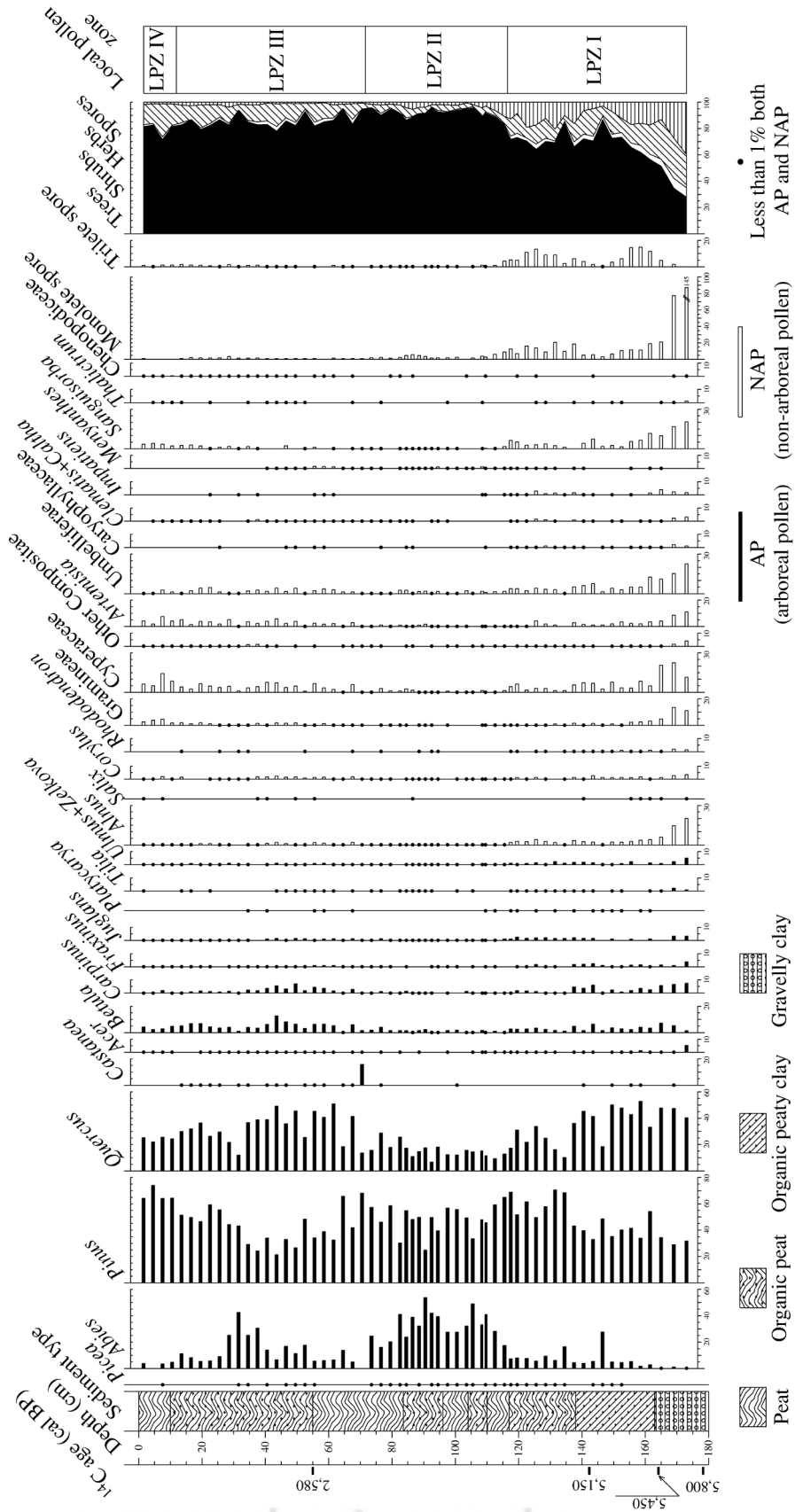


Fig. 5. Pollen taxa percentage diagram (left, resolved diagram) and total pollen percentage diagram (right, cumulative diagram) of selected taxa for sediment collected from the Yongneup moor at Mt. Daamsan, Gangwon-do province, Korea. The percentage calculations are based on the total arboreal pollen sum (pollen percentage diagram) and total pollen and spore sum (total pollen percentage diagram). Cal BP, calibrated years before present.

LPZ-I (*Quercus-Pinus* zone)

The pollen spectra during the period 5,900-4,800 cal BP (lower sediment) were characterized by high percentages of *Quercus* (30-50%) and *Pinus* (29-70%), whereas the frequency of *Abies* was < 10% throughout the LPZ-I zone. Among tree pollen taxa, *Betula* and *Carpinus* revealed relatively high percentages.

Pollen of the shrubby *Alnus*, an indicator of the local environment, was relatively abundant. This taxon appeared in approximately 20% of the bottom layers and then decreased gradually toward the upper layer of LPZ I. Herbaceous pollen taxa such as Cyperaceae, *Artemisia*, Umbelliferae, and *Sanguisorba* predominated in this pollen zone compared to other herbaceous taxa. In contrast, the pollen assemblages of *Abies* and *Pinus* began to increase and that of *Quercus* began to decrease in this pollen zone.

LPZ-II (*Pinus-Abies-Quercus* zone)

The LPZ-II zone was estimated to be 4,800-3,400 cal BP and was characterized by rapid increases in *Abies* and *Pinus* and by abrupt decreases in *Quercus*, including other deciduous broad-leaved taxa. The percentage *Abies* pollen generally exceeded 15% throughout this zone and occasionally reached 49%. *Pinus* pollen occupied 24-65%.

Compared with the pollen frequency in the LPZ-I zone, *Quercus* in the LPZ-II zone showed lower percentages of 10-17%. Additionally, deciduous broad-leaved tree pollen taxa composed of a tall-tree layer comprised < 1%. Herbaceous pollen and spores occurred in relatively low percentages compared with the LPZ-I zone. *Quercus* began to increase again and *Abies* and *Pinus* began to decrease in this pollen zone.

LPZ-III (*Quercus-Pinus-Abies* zone)

The LPZ-III zone was estimated to be 3,400-390 cal BP. The vegetation information on the LPZ-III zone was obtained from the 72-12 cm depth sediment layer. The LPZ-III zone was characterized by an increase in *Quercus* and a decrease in *Pinus* and *Abies*. *Quercus* pollen exceeded 50%, whereas that of *Pinus* occurred at 68% in the lower layer, but decreased to 21% in the upper layer of the LPZ-III zone. *Abies* pollen occurred at an overall frequency of 10-15%. Among the other AP pollens, *Betula* and *Carpinus* occurred in relatively high percentages. Herbaceous pollen and spores were also slightly higher than those in the LPZ-II zone.

LPZ-IV (*Pinus-Quercus* zone)

The LPZ-IV zone was estimated to be 390-present cal BP was the uppermost zone including sediment at depths of 12-0 cm. This zone was characterized by rapid increases in *Pinus* and by abrupt decreases in *Abies*.

Abies pollen, which occurred at relatively high percentages in the LPZ-III zone decreased to < 10% and, thereafter, this taxon showed a continually decreasing trend in this zone. In contrast, *Pinus* pollen showed a rapidly increasing trend to 74%. The *Quercus* pollen assemblage was consistently maintained at a frequency of 22-25%, although the frequency was lower than that in the LPZ-III zone.

DISCUSSION

From the ¹⁴C-dating results, the bottom sediments began to form in the Yongneup moor approximately 5,900 cal BP. The successive pollen spectra in Yongneup moor sediment provided information for reconstructing the changes in past vegetation on local and regional scales in the mid-late Holocene. The Korean peninsula vegetation types for reconstructing past vegetation were based on Yim and Kira (1975), Yim (1977a, 1977b), and Kim (1992). Decisions about past vegetation based on the pollen assemblages were based on Choi (1997, 1998) and Song (2002). Chang et al. (1987) has conducted analytical pollen studies on this moor, but they did not perform radiocarbon dating. Furthermore, their RPF calculation method differed from that used in the present study. Therefore, we compared our relative similarities with those of the main AP RPFs, as a direct comparison to their results was difficult.

According to the pollen assemblage spectra in the LPZ-I zone, which occurred 5,900-4,800 cal BP, it was inferred that past vegetation along the foot to the slope of Mt. Daeamsan was composed predominantly of *Q. mongolica* mixed forest with *Carpinus*, *Betula*, and *Fraxinus*. In other words, cool temperate central/montane deciduous broad-leaved forest probably developed in this area. However, this suggested that the cool temperate northern/alti-montane mixed coniferous and deciduous broad-leaved forest developed at the summit and the surrounding area of Mt. Daeamsan. The past forest on the top of the mountain was dominated by *Q. mongolica*, *P. koraiensis*, *Abies holophylla*, and *A. nephrolepis*.

The *Alnus* pollen assemblage and herbaceous taxa in the LPZ-I zone had noticeably higher percentages in contrast to the upper layers, indicating that a depressive bottom during the initial period of Yongneup develop-

ment maintained poor drainage conditions originating from the formation of wet soil including fine sand, silt, and clay. *Alnus japonica* is an indicator shrub that grows on wet soil such as moors, fens, and swamps. The abundance of *Alnus* pollen in the LPZ-I zone indicated that *A. japonica* inhabited Yongneup bog or its surrounding mountain foot. In other words, other AP trees did not develop in the moor and/or nearby mountain areas due to the wet soil conditions. In contrast, herbaceous taxa such as Cyperaceae, *Artemisia*, Umbelliferae, and *Sanguisorba* occurred at relatively higher densities.

Considering the variations in fossil pollen and spores during LPZ-II (4,800-3,400 cal BP), it is estimated that subalpine coniferous forest developed at the outskirts of Yongneup moor and along the ridge to the summit of Mt. Daeamsan. This indicates that the subalpine coniferous forest, composed of *A. holophylla*, *A. nephrolepis*, and *P. koraiensis*, spread quickly along the foot up to the summit of Mt. Daeamsan and occupied the surrounding area during this period.

The high percentages of subalpine pollen taxa in the LPZ-II occurred because the subalpine coniferous forest, such as *A. holophylla*, *A. nephrolepis*, and *Picea jezoensis* descended vertically and expanded at lower elevations. Additionally, it can be inferred that the cool temperate northern/alti-montane mixed coniferous and deciduous broad-leaved forest, composed of *Q. mongolica* and *P. koraiensis*, covered most of the mountainous parts below the Mt. Daeamsan piedmont. Judging from the fluctuations in *Pinus* and *Quercus*, this pollen zone was comparable to pollen zone I of Chang et al. (1987); however, *Abies* did not appear clearly as in our study.

During the period from 3,400-390 cal BP (LPZ-III), *Abies* and *Pinus* taxa, which were distributed during the previous LPZ-II period, retreated to the summit or areas further north than Mt. Daeamsan. Therefore, it is inferred that the retreated *Abies* and *Pinus* taxa were replaced by *Quercus* and other deciduous broad-leaved species. In other words, *Quercus*, which dominated below the mountainside in the previous period, developed above the mountainside or near the summit areas, indicating that cool temperate northern/alti-montane mixed coniferous and deciduous broad-leaved forest covered most of the summit area, whereas cool temperate central/montane deciduous broad-leaved forest covered most of the lower regions. Judging from the fluctuations in *Pinus* and *Quercus*, this pollen zone was comparable to pollen zone II of Chang et al. (1987).

The reconstructed vegetation types of this period nearly coincided with those of LPZ-I. However, the per-

centages of *Abies* and *Pinus* were higher than those of LPZ-I until 1,300 cal BP. These results reflect that the frequency of evergreen subalpine coniferous species, such as *A. holophylla*, *A. nephrolepis*, and *P. koraiensis*, were relatively higher than that of the *Quercus-Pinus* zone (LPZ-I), which was the bottom sediment. Accordingly, it is inferred that the subalpine coniferous forest ascertained in the previous *Pinus-Abies-Quercus* zone (LPZ-II) also developed at elevations close to the Mt. Daeamsan summit during the lower to middle parts of LPZ-III.

In LPZ IV (390-present cal BP), the most conspicuous pollen frequency feature was a sharp increase in *Pinus* and a distinct decrease in *Abies*. While *Quercus* was less frequent than the previous period, *Quercus* pollen was maintained continuously within a range of 22-25%. Such results indicate that coniferous *Abies* taxa, such as *A. holophylla* and *A. nephrolepis*, occurred individually below the middle region and at the community level above Mt. Daeamsan during the previous LPZ-III pollen zone and were rapidly replaced by *Pinus* (*P. densiflora*) during the LPZ-IV period. This pollen zone was comparable to pollen zone III of Chang et al. (1987).

Based on the pollen composition, the vegetation type surrounding Mt. Daeamsan during LPZ IV was probably cool temperate central/montane mixed deciduous broad-leaved and coniferous forest. The abrupt increase in *Pinus* pollen in this pollen zone may be pollen dispersed from secondary *P. densiflora* forests formed by humans, who were establishing in the lower regions around Mt. Daeamsan.

CONCLUSION

A pollen analysis and radiocarbon dating of sediment peat collected from Yongneup moor were conducted to reconstruct past vegetation in the mountainous region of the central Korean peninsula during the Holocene. The 61 pollen analysis samples and the five ¹⁴C-dating samples were obtained from depths of 180 cm to the surface. Of the samples collected, the sediments from depths of 130 cm to the surface were sequential peat.

The year of the bottom sediment was approximately 5,900 cal BP. Four LPZs were identified from the pollen analysis results at Yongneup moor. The past vegetation inferred from the pollen record included: (1) cool temperate central/montane deciduous broad-leaved forest (LPZ-I) during the period from 5,900-4,800 cal BP; (2) cool temperate northern/alti-montane mixed coniferous and deciduous broad-leaved forest (LPZ-II) during the period

from 4,800-3,400 cal BP, (3) cool temperate central/montane deciduous broad-leaved forest (LPZ-III) during the period of 3,400-390 cal BP and (4) cool temperate central/montane mixed deciduous broad-leaved and coniferous forest (LPZ-IV) during the period from 390-present cal BP. These pollen zones represent the vegetation dynamics at the different levels and reveal several original features of the areas around Yongneup bog and Mt. Daeamsan since the Holocene.

Although the pollen analysis technique is useful to reconstruct past vegetation, it should be applied in parallel with ¹⁴C-dated sequences of other subalpine wetlands such as bogs and fens to elucidate past vegetation on the Korean peninsula. Additionally, insufficient data are available, except for Yongneup bog, so further efforts should be made to conduct additional palynological analyses.

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